



Comparison of inference methods for estimating semivariogram model parameters and their uncertainty: The case of small data sets

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ARTICLE INFO

Article history:

Received 2 April 2012

Received in revised form

29 May 2012

Accepted 1 June 2012

Available online 12 June 2012

Keywords:

Maximum likelihood

Cross-validation

Least squares

Chi-squared field

Root mean square error

ABSTRACT

The semivariogram model is the fundamental component in all geostatistical applications and its inference is an issue of significant practical interest. The semivariogram model is defined by a mathematical function, the parameters of which are usually estimated from the experimental data. There are important application areas in which small data sets are the norm; rainfall estimation from rain gauge data and transmissivity estimation from pumping test data are two examples from, respectively, surface and subsurface hydrology. Thus a benchmark problem in geostatistics is deciding on the most appropriate method for the inference of the semivariogram model.

The various methods for semivariogram inference can be classified as indirect methods, in which there is an intermediate step of calculating the experimental semivariogram, and direct approaches that obtain the model parameter values directly as the values that minimize some objective function.

To avoid subjectivity in fitting models to experimental semivariograms, ordinary least squares (OLS), weighted least squares (WLS) and generalized least squares (GLS) are often used. Uncertainty evaluation in indirect methods is done using computationally intensive resampling procedures such as the bootstrap method.

Direct methods include parametric methods, such as maximum likelihood (ML) and maximum likelihood cross-validation (MLCV), and non-parametric methods, such as minimization of cross-validation statistics (CV).

The bases for comparing the previous methods are the sampling distribution of the various parameters and the “goodness” of the uncertainty evaluation in a sense that we define. The final questions to be answered are (1) which is the best method for estimating each of the semivariogram parameters? (2) Which is the best method for assessing the uncertainty of each of the parameters? (3) Which method best selects the functional form of the semivariogram from among a set of options? and (4) which is the best method that jointly addresses all the previous questions?

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1. Introduction

Since the early applications of geostatistics in mining (Matheron, 1963), the range of applications has been extended significantly to include hydrogeology (e.g. Kitanidis, 1997), petroleum geology (e.g. Deutsch, 2002; Caers, 2005), meteorology and climatology (e.g. Diodato and Ceccarelli, 2005), soil science (e.g. Goovaerts, 1997), remote sensing (e.g. Curran and Atkinson, 1998), geographical information systems (e.g. Burrough, 2001), image analysis (e.g. Chica-Olmo and Abarca-Hernández, 2000), ecology (e.g. Liebhold et al., 1993), econometrics (e.g. Griffith and

Paelinck, 2011), health sciences (e.g. Kelsall and Wakefield, 2002), and many other disciplines that use geostatistics in the analysis of spatial (and temporal) data. Geostatistics is applied to problems that require, inter alia, the quantification of spatial variability, optimal interpolation, scenario generation for risk analysis, and sampling design. With the exception of recent advances in multiple-point geostatistics (e.g. Strebelle, 2002), the basis of all geostatistical applications is a semivariogram model that describes the spatial variability of a random function that models the spatial (or regionalized) variable of interest. The semivariogram provides a means of detecting the defining characteristics of the spatial variability of a spatial variable such as anisotropic variability for different spatial directions (geometric and zonal anisotropies), different scales of spatial variability (nested structures), scales of variability shorter than the shortest sampling distance (nugget variance), measurement errors (nugget

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List of acronyms

RF	random function	WLS2	weighted least squares with binned semivariogram data
OLS	ordinary least squares with the complete semivariogram cloud	GLS	generalized least squares with binned semivariogram data
WLS	weighted least squares, (using the semivariogram cloud for distances less than 30 units); this is done by applying zero weights to the semivariogram cloud for distances greater than 30 units	ML	maximum likelihood
WLS1	weighted least squares with the semivariogram cloud for distances less than a distance threshold	CV1	cross-validation
		CV2	cross-validation with orthonormal residuals
		MLCV1	maximum likelihood cross-validation
		MLCV2	maximum likelihood cross-validation with orthonormal residuals
		RMSE	root mean square error

variance), differentiability of the variable (behavior of the semivariogram close to the origin), and cyclic spatial patterns (semivariogram hole effect). In this sense the semivariogram is important *per se* and is an essential tool of spatial statistics; in addition, a semivariogram model is required for kriging, cokriging, simulation and optimal design methodologies. Generally, in practice, the semivariogram model is unknown and is estimated from the experimental data. The estimation of the semivariogram is a process of statistical inference that has generated a significant amount of scientific literature covering many aspects of the problem. The following problems are of significant interest in semivariogram parameter inference:

- (1) Robust semivariogram estimation (e.g. Dowd, 1983; Genton, 1998).
- (2) Non-parametric semivariogram estimation (e.g. Sampson and Guttorp, 1992).
- (3) Estimation of the trend or drift. The drift is important *per se* and special procedures may be used (e.g. Visser et al., 2009) or it may be implicitly incorporated in estimation by using a moving window for sample selection (Journel and Rossi, 1989).
- (4) Optimal sampling design for semivariogram estimation (e.g. Bogaert and Russo, 1999).
- (5) Bayesian estimation or inclusion of *a priori* knowledge about the semivariogram parameters coded in an *a priori* probability density function (e.g. Pardo-Igúzquiza, 1999a).
- (6) The impact of semivariogram estimation errors on kriging interpolation (e.g. Zimmerman and Zimmerman, 1991).
- (7) Transformation of the data (normal scoring, Box-Cox transform, etc) (e.g. Gringarten and Deutsch, 2001).
- (8) Criteria for model selection (e.g. Ye et al., 2008).
- (9) Re-parameterization of semivariogram parameters to achieve parameter orthogonality in maximum likelihood estimation (e.g. Diggle and Ribeiro, 2007).

However, in order to be specific we limit our study to the estimation of the basic Matheron representation of the semivariogram (Matheron, 1963), that is, a second order stationary random function (i.e., constant drift) with an isotropic semivariogram model and allowing for a nugget effect. Thus, in general, there are three semivariogram parameters of interest: range, nugget variance and partial sill. The variance, or total sill, is the sum of nugget variance and partial sill. If there is no nugget variance then the two parameters of interest are the range and the sill of the semivariogram (variance of the random function). There is a subtle difference between explicitly excluding a nugget variance from the model and estimating the nugget variance to be zero:

- It may be concluded in advance that there is no nugget variance and thus implicitly the nugget variance has a value of zero with no uncertainty.

- The nugget variance may be explicitly estimated. Even if the estimated value is zero, the estimate will have an associated uncertainty and the model complexity is increased by one parameter with the consequence of increasing the complexity of the estimation procedure.

This basic Matheron model representation is adequate for the case of small data sets that we consider in this work. For large data sets, such as remotely sensed data, there is sufficient information in the data to provide acceptable semivariogram parameter estimates using the standard inference methods. However, very small data sets do not convey sufficient information for the inference of complex spatial variability models. For example, there may be too few data from which to estimate directional statistical variability and hence assess anisotropy. In such cases the most reliable strategy is to choose a simple model with the minimum number of parameters. The small data set problem is important in applications because there are many areas in geosciences in which data are few; for example, because data collection is expensive (e.g., rain gauges in remote regions or the cost of drilling a borehole) or because the measurement modifies the variable that is being measured (e.g., for potential underground hazardous waste repositories in which a borehole alters the hydraulic connectivity of the rock).

In geoscience applications, a small sample is generally defined as one with a few tens of values, such as between 10 and 100 (O'Brien and Griffiths, 1965). In fact, in small sampling theory (Spiegel and Stephens, 2008) a small sample is considered to be one with less than 30 values. However, this refers to applications in classical statistics where the data are considered as realizations of independent and identically distributed random variables. When spatial correlation is present, the effective sample size may be smaller than the actual number of sample values and for this reason we consider small sample sizes to be those with up to 100 data. Typical examples of small sample sizes in geostatistical studies include sample sizes of 36 and 100 (Mardia and Marshall, 1984), 25, 50, and 100 (Krajewski and Duffy, 1988), 16 and 36 (Zimmerman and Zimmerman, 1991), 32 and 72 (Russo and Jury, 1987), and 12 (Sampson and Guttorp, 1992).

An additional issue when estimating the semivariogram parameters from a small sample is that there is an unavoidable uncertainty associated with them and that uncertainty must be assessed and assigned to the estimates. This should be the norm, rather than the exception, in acceptable statistical practice (Bard, 1974). The next section describes the inference methods, the methods of uncertainty assessment, the design of the experiments and the criteria for the comparisons.

2. Methodology

In this section we provide a brief review of the methods for inference and uncertainty evaluation.

2.1. Inference methods

The indirect methods involve two steps: firstly, the calculation of an experimental semivariogram and, secondly, fitting a theoretical model to the experimental semivariogram. The semivariogram model is defined as a functional form (e.g. spherical, exponential, Gaussian, Matérn) and a set of parameters (range, nugget variance and partial sill). The standard way of calculating the experimental semivariogram for irregularly located data is to use a binning process in which the bins are defined by a lag value plus/minus a lag tolerance and the squared differences of all data pairs are assigned to the bins (e.g., Chilès and Delfiner, 2012). As the binning process leads to a loss of information, a preferable alternative is to calculate the semivariogram cloud. Given n experimental data, the semivariogram cloud is the set of $n(n-1)/2$ semivariogram values calculated as the squared difference of each pair of experimental data:

$$\gamma(h_{ij}) = \frac{1}{2} [z(\mathbf{u}_i) - z(\mathbf{u}_j)]^2, \quad \{i = 1, \dots, n-1; j = i+1, \dots, n\} \quad (1)$$

where $\{z(\mathbf{u}_i), z(\mathbf{u}_j)\}$ is a pair of experimental data. \mathbf{u}_i are the spatial coordinates of the i th datum. For the purposes of this paper we use data measured in two-dimensional space, i.e., $\mathbf{u}_i = \{x_i, y_i\}$. The method of least squares is an objective method for fitting a theoretical semivariogram model to an experimental semivariogram (Cressie, 1985). The semivariogram parameter estimates are those that minimize the difference between the theoretical model and the experimental semivariogram:

$$\hat{\theta} = \arg \min_{\theta} (\hat{\mathbf{G}} - \mathbf{G}(\theta))^T \Sigma^{-1} (\hat{\mathbf{G}} - \mathbf{G}(\theta)) \quad (2)$$

where θ is the $p \times 1$ vector of semivariogram parameters (range, nugget variance and partial sill); $\hat{\theta}$ is the $p \times 1$ vector of semivariogram parameter estimates or arguments of the minimum, that is, the values that minimize the objective function $(\hat{\mathbf{G}} - \mathbf{G}(\theta))^T \Sigma^{-1} (\hat{\mathbf{G}} - \mathbf{G}(\theta))$; $\hat{\mathbf{G}}$ is the $(n(n-1)/2) \times 1$ vector of the experimental semivariogram values, that is, the semivariogram cloud; $\mathbf{G}(\theta)$ is the $(n(n-1)/2) \times 1$ vector of the theoretical semivariogram values for the same distances as for the semivariogram cloud. These values depend on θ , the vector of semivariogram parameters; Σ is the $(n(n-1)/2) \times (n(n-1)/2)$ matrix in which the elements are defined by the least squares method used, i.e. ordinary least squares (OLS), weighted least squares (WLS) or generalized least squares (GLS); and the superscript T denotes a transpose vector or transpose matrix and the superscript -1 indicates inverse matrix.

In OLS the matrix Σ is a diagonal matrix with a value of 1 on the diagonal $\Sigma = \text{diag}(1, 1, \dots, 1)$.

In WLS the matrix Σ is a diagonal matrix with variance values on the diagonal $\Sigma = \text{diag}(\sigma_1, \sigma_2, \dots, \sigma_m)$, with $m = n(n-1)/2$. Where $\sigma_k = \text{var}[\gamma(h_{ij})]$ is the variance of the experimental semivariogram (as defined in (1)) for the pair of data $\{z(\mathbf{u}_i), z(\mathbf{u}_j)\}$ that occupy the k th location in the sequence of m diagonal elements. A simplification of σ_k , to avoid the calculation of $\text{var}[\gamma(h_{ij})]$, is to use an empirical function, such as the square of the distance h_{ij}^2 or the square of the semivariogram value $\gamma(h_{ij}; \theta)$; see, for example, Pardo-Igúzquiza (1999b). There are two drawbacks of WLS fitting, when used with binned data: first, the results depend on the parameters used for binning (e.g., number of directions, angle tolerance, number of lags, and lag tolerance) and second, the results are biased (Diggle and Ribeiro, 2007).

In GLS the elements of the matrix Σ are the covariances between the components of the semivariogram cloud. The general element is

$$[\Sigma]_{uv} = \text{cov}[\gamma(h_{ij}), \gamma(h_{k,l})] \quad (3)$$

The variance-covariance matrix of the semivariogram cloud can be calculated analytically by assuming that the data follow a multivariate Gaussian distribution (Pardo-Igúzquiza and Dowd, 2001). More generally, without making the Gaussian assumption, a bootstrap procedure can be used (Olea and Pardo-Igúzquiza, 2011). These formulations are equally valid for the semivariogram cloud and for the traditional calculation using the binning procedure.

For the direct methods, the semivariogram parameters are obtained directly without the intermediate step of calculating an experimental semivariogram. In turn, this group of inference methods can be divided into parametric and non-parametric methods. The most important of the parametric methods is the method of maximum likelihood (ML). The ML estimates of the semivariogram model parameters are those that minimize the negative log-likelihood function (Kitanidis and Lane, 1985; Mardia and Marshall, 1984):

$$L(\theta | \mathbf{z}) = \frac{n}{2} \ln(2\pi) + \frac{1}{2} \ln |\mathbf{C}| + \frac{1}{2} (\mathbf{z} - \mathbf{m})^T \mathbf{C}^{-1} (\mathbf{z} - \mathbf{m}) \quad (4)$$

where \mathbf{z} is the $n \times 1$ vector of the experimental data; \mathbf{C} is the $n \times n$ covariance matrix for the experimental data. The general term is $[\mathbf{C}]_{ij} = \text{cov}(z(\mathbf{u}_i), z(\mathbf{u}_j)) = \sigma_z^2 - \gamma(h_{ij})$; \mathbf{m} is the $n \times 1$ vector of the identical elements equal to the global mean of \mathbf{Z} : m_z ; σ_z^2 is the variance of the random function \mathbf{Z} and equal to the semivariogram sill.

Another parametric method is ML cross-validation (Samper and Neuman, 1989) which is a special case of cross-validation as shown below (Kitanidis, 1991).

A non-parametric method for estimating the semivariogram model parameters is cross-validation (Lebel and Bastin, 1985). Given the n experimental data, \mathbf{z} , the k th datum is omitted and the rest of the data are used to estimate $z(\mathbf{u}_i)$ by any linear unbiased procedure to yield the estimate:

$$\hat{z}(\mathbf{u}_k) = \sum_{\substack{i=1 \\ i \neq k}}^n \lambda_i z(\mathbf{u}_i), \quad (5)$$

with

$$\sum_{\substack{i=1 \\ i \neq k}}^n \lambda_i = 1. \quad (6)$$

The estimation variance, s_k^2 , can be expressed as a function of the semivariogram model parameters:

$$s_k^2 = 2 \sum_{\substack{i=1 \\ i \neq k}}^n \lambda_i \gamma(h_{ik}) + \sum_{\substack{i=1 \\ i \neq k}}^n \sum_{\substack{j=1 \\ j \neq k}}^n \lambda_i \lambda_j \gamma(h_{ij}) \quad (7)$$

The standardized estimation error is given by:

$$e_k = \frac{\hat{z}(\mathbf{u}_k) - z(\mathbf{u}_k)}{s_k} \quad (8)$$

and the following statistic can be calculated:

$$S_2 = \frac{1}{n} \sum_{i=1}^n e_i^2 \quad (9)$$

If the data are consistent with the semivariogram model, S_2 should be one. Thus the semivariogram parameter vector can be estimated by cross-validation as the solution to the equation:

$$S_2 = 1 \quad (10)$$

However, Kitanidis (1991) has shown that there are difficulties with the standardized residuals and it is better to use orthonormal residuals ε_k as defined in Kitanidis (1991). The equivalent

statistic to S_2 with orthonormal residuals is Q_2 ,

$$Q_2 = \frac{1}{n-1} \sum_{i=1}^n \varepsilon_k^2 \tag{11}$$

and the semivariogram parameter vector can be estimated by cross-validation as the solution to the equation: $Q_2=1$, which is equivalent to Eq. (10) but with orthonormal residuals instead of standard residuals. However, the solution of $Q_2=1$, is unique only if there is a single semivariogram parameter to be estimated. For a larger number of parameters, there are several vectors θ that satisfy Eq. (11). Thus a measure of overall accuracy is introduced as, for example, (Kitanidis, 1991):

$$\frac{1}{n} \sum_{i=1}^n \ln(s_i^2) \tag{12}$$

The estimates of the parameter vector θ are the values that minimize Eq. (12) subject to the constraint given in Eq. (11). Kitanidis (1983, 1991) shows that this solution is equivalent to the ML method of minimizing the negative Gaussian log-likelihood:

$$\arg \min_{\theta} \left\{ \sum_{i=1}^n \ln(s_i^2) + \sum_{i=1}^n \varepsilon_i^2 \right\} \tag{13}$$

The latter method is similar to the maximum likelihood cross-validation (MLCV) of Samper and Neuman (1989).

2.2. Methods of uncertainty assessment

For least squares estimation, the approximate variance-covariance matrix of the estimates, under the assumption that their distribution is Gaussian, is given by the matrix C_{θ} (Menke, 1984):

$$C_{\theta} = [J^T \Sigma^{-1} J]^{-1} \tag{14}$$

where J is the $K \times p$ Jacobian with ij th element defined by $[J]_{kj} = \partial \gamma(h_k) / \partial \theta_j$, evaluated at the least squares estimates. K is the number of elements of the semivariogram cloud and p is the number of semivariogram model parameters.

Another possibility for evaluating uncertainty is by a computationally intensive method such as the bootstrap. We use the procedure of Solow (1985) to generate bootstrap samples from independent and identically distributed data after a transformation of the original data. This procedure is preferable to bootstrapping the semivariogram as shown in Olea and Pardo-Igúzquiza (2011).

For each bootstrap sample, a set of parameters θ is estimated, providing an estimate of the sampling distribution of the parameters from which joint measures of uncertainty may be obtained (Olea and Pardo-Igúzquiza, 2011).

One advantage of a parametric method, such as ML, is that it provides three ways of evaluating uncertainty: the Fisher information matrix, likelihood regions and confidence intervals from the likelihood ratio test statistic.

The Fisher information matrix $F(\theta)$ is given by the expression (Kitanidis, 1983; Mardia and Marshall, 1984):

$$[F(\theta)]_{ij} = \frac{1}{2} \text{Tr} [C^{-1} C_i C^{-1} C_j], \tag{15}$$

where $\text{Tr}[\cdot]$ denotes the trace of a matrix, and

$$C_i = \frac{\partial C}{\partial \theta_i} \tag{16}$$

The inverse of the Fisher information matrix is an approximation to the covariance matrix of the semivariogram parameter

estimates:

$$C_{\theta} = [F(\theta)]^{-1} \tag{17}$$

The likelihood ratio statistic is defined as (Kalbfleisch, 1979):

$$D(\theta) = 2[L(\theta) - L(\theta_{ML})] \tag{18}$$

For large n , $D(\theta)$ is approximately distributed as a chi-squared distribution with p degrees of freedom (McCullagh and Nelder, 1989). For small n , this approximation is often more accurate than the Fisher information matrix (Kalbfleisch, 1979). Using this approximation it is possible to construct confidence regions for semivariogram parameters (Pardo-Igúzquiza and Dowd, 2003). A detailed description is given in Pardo-Igúzquiza et al. (2009, p. 29).

For uncertainty measurements of the semivariogram parameters estimated by cross-validation and orthonormal residuals, the confidence regions can be calculated from the duality between confidence regions and hypothesis tests (Rice, 1995). Kitanidis (1991) shows that under the Gaussian assumption for the orthonormal residuals, the statistic:

$$(n-1)Q_2, \tag{19}$$

follows a chi-squared distribution with $n-1$ degrees of freedom. Thus, if the acceptance region of a test at a significance level α is $A(Q_2)$, then the set:

$$R(Q_2) = \{\theta : Q_2 \in A(Q_2)\}, \tag{20}$$

is a $100(1-\alpha)\%$ confidence region for θ .

Another possibility is to use the Fisher information matrix, Eq. (15), together with the full formulation of maximum likelihood cross-validation and the assumption of Gaussian residuals.

2.3. Design of the experiments for the comparison

Fig. 1 shows the design of the experiment for comparing methods for inferring semivariogram parameters and the associated assessment of uncertainty. The underlying random function (RF) model is a zero-mean, second-order, stationary function with a given semivariogram model defined by a functional form (spherical, exponential or Gaussian) and the semivariogram model parameters θ . A realization of the RF is generated using

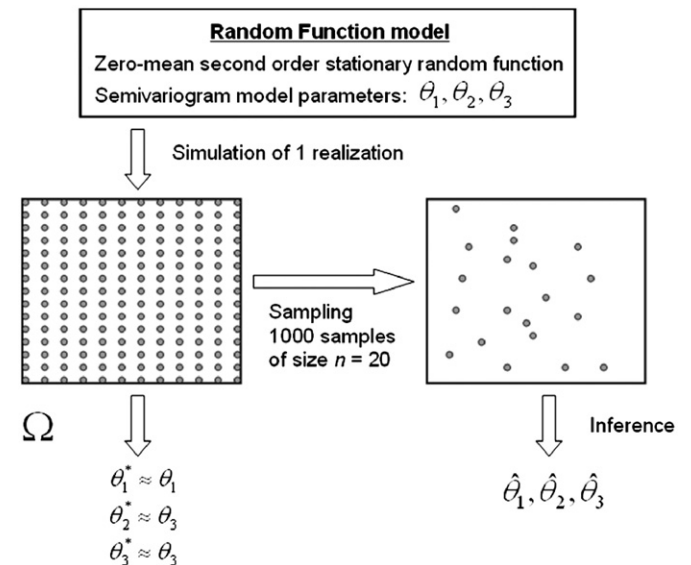


Fig. 1. Design of the experiment for selecting small sample sets on which to compare estimators.

any of the many methods available for non-conditional geostatistical simulation (Chilès and Delfiner, 2012). Because the simulation is generated on a grid defined on a finite region Ω (Fig. 1), there is an ergodic fluctuation in the simulation (Deutsch and Journel, 1998) in the sense that, if the parameters are estimated from the simulation, the semivariogram parameters θ^* of a realization, may differ from the underlying theoretical values θ . If many realizations of the RF are generated, then for any sound geostatistical simulation method:

$$E\{\theta^*\} = \theta \quad (21)$$

For the purposes of this paper a single realization is sufficient. The complete realization is a reality on the scale of the simulation grid and the complete realization is such that

$$\theta^* \approx \theta \quad (22)$$

We can now sample this realization using the number of samples that typifies a small data set. The problem is to infer the underlying parameters θ from the small sample set. To do so, we are seeking efficient inference methods that are able to give estimates close to the underlying values θ and provide efficient measures of uncertainty in a sense that we define later.

The factors that have the most significant influence on the outcomes of the experiment are:

- The sample size.
- Whether or not the RF is Gaussian.
- The nugget/variance ratio.
- The ratio of the practical range to the characteristic length of the study area. The characteristic length of the study area is a measure of the size of the simulation region Ω . If this region is a square, the characteristic length is the length of the side of the square.

We use four simulated fields:

- (A.) Gaussian field with no nugget variance, shown in Fig. 2A.
- (B.) Gaussian field with nugget variance equal to 50% of the total variance, shown in Fig. 2B.
- (C.) Non-Gaussian random field with no nugget variance, shown in Fig. 2C.
- (D.) Non-Gaussian random field with nugget variance equal to 50% of the total variance, shown in Fig. 2D.

Each field is simulated on a 101×101 grid of points with unit distance between points along the X and Y axes.

For each of the four fields we consider practical range/characteristic length ratios of 0.2 (Small), 0.4 (Medium) and 0.6 (Large). This gives twelve cases (four fields and three ratios for each) and for each case we use sample sizes of 10, 30, 50, 70 and 90. For each sample size we sample the field 1000 times to give 1000 data sets. Although we do not use all these data files, we provide them as a benchmark data set for further comparisons by others. The files are available at <http://www.iamg.org/CGEditor/index.htm>. The semivariogram model is exponential with variance 10 units and range 6 units (practical range 18 units of length). In cases B and D the nugget variance is 5, that is 50% of the total variance. The particular values of 6, 10 and 5 have no influence on the results.

The non-Gaussian RF is a chi-squared random field with one degree of freedom (Pardo-Igúzquiza and Dowd, 2005). Although the chi-squared RF is obtained by squaring a Gaussian RF, it is not possible to recover the original Gaussian field by a transformation of the chi-squared RF, and in this sense, this represents a difficult case and a very good example of non-Gaussian RF.

2.4. Criteria for the evaluations

Sampling consists of selecting the locations of the n data for each sample at random from the grid locations of the complete realization (Fig. 1). For each sample, each inference procedure provides a set of estimates $\hat{\theta}$ and associated uncertainty measures $U(\hat{\theta})$, which, for the general case, is the variance-covariance matrix of the estimates $U(\hat{\theta}) = C_{\hat{\theta}}$ or a confidence region such as that given in Eq. (19). Provided a distribution is assumed for the estimates, the variance-covariance matrix can be used to construct confidence regions for the set of estimates. We are interested in comparing two performance measures for the inference methods: how well they estimate the underlying semivariogram parameters θ and the ability of the uncertainty measures to evaluate the true uncertainty.

2.4.1. Methods of evaluating estimators

The most common measure of the performance of the estimators is the mean square error (MSE) defined as

$$\text{MSE}(\hat{\theta}) = \mathbf{b}\mathbf{b}^T + \text{var}(\hat{\theta}) \quad (23)$$

i.e., the square of the bias plus the variance, where the bias is given by

$$\mathbf{b} = \bar{\theta} - \theta \quad (24)$$

$$\bar{\theta} = E\{\hat{\theta}\} \quad (25)$$

and the variance is

$$\text{var}(\hat{\theta}) = E\{(\hat{\theta} - \bar{\theta})(\hat{\theta} - \bar{\theta})^T\} \quad (26)$$

The MSE provides a measure of the closeness of the estimates to the true underlying parameters. The square root of the mean square error (RMSE) is also used. As the sampling distribution of the estimator is obtained for each parameter (Fig. 1), other measures of performance could also be obtained.

2.4.2. Methods of evaluating uncertainty measures

An uncertainty measured is assigned to each estimate. The larger the uncertainty the less will be the reliability of the estimate. One way of comparing uncertainty measures is by using interval estimates constructed from the uncertainty measures. The two measures for the intervals are size and coverage probability. The ideal is the narrowest interval with the maximum possible coverage. There is a trade-off between size and coverage because, usually, larger intervals have larger coverage. However, very large intervals, even with large coverage, are of little use; for example, for a single parameter, the interval $(-\infty, +\infty)$ has a coverage of one, but it is obviously meaningless. For a single parameter, the size of the confidence set is a length but, in the general case, it is a volume of multiple parameters. As there are many confidence sets with the same probability coverage, we use the narrowest interval with the highest probability coverage. In addition, the estimate must be inside the confidence set.

3. Results

For an underlying Gaussian random function, Pardo-Igúzquiza (1998) shows that ML performs better than OLS, WLS or MLCV. Although we note that GLS and CV with orthonormal residuals were not used in this comparison.

For the results reported here we used 100 samples for each estimator although we have provided the files with 1000 data sets for each of the 40 cases. We show the results for the worst case scenario. That is, the underlying RF model is non-Gaussian and there is no possibility of transforming the data prior to applying

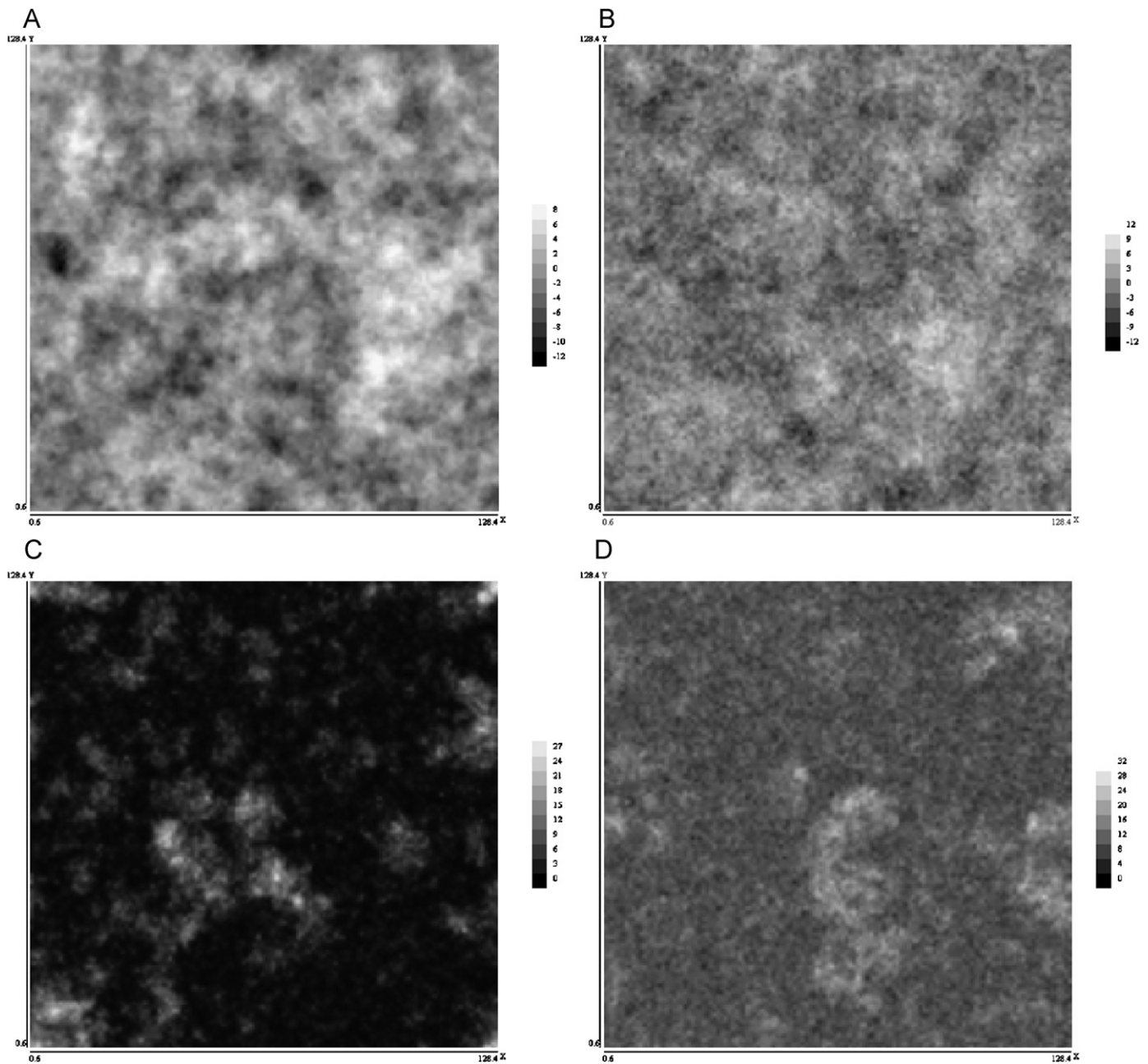


Fig. 2. (A) Realization of a zero-mean Gaussian field with exponential semivariogram with range 6 units (practical range 18 units of distance). The nugget variance is zero and the variance is 10. (B) Realization of a zero-mean Gaussian field with exponential semivariogram with range 6 units (practical range 18 units of distance). The nugget variance is 50% of the total variance. The total variance is 10, the nugget variance is 5 and the partial sill is 5. (C) Realization of a chi-square field with exponential semivariogram with range 6 units (practical range 18 units of distance). The nugget variance is zero and the total variance is 10. (D) Realization of a chi-square field with exponential semivariogram with range 6 units (practical range 18 units of distance). The nugget variance is 50% of the total variance. The variance is 10, the nugget variance is 5 and the partial sill is 5.

the estimator or of correcting the estimates. We use a practical range/field length ratio of 0.4 and compare the estimators (as given in the list of acronyms).

4. Discussion

4.1. Estimation of semivariogram model parameters

4.1.1. Non-Gaussian RF with no nugget variance

In this case there are two parameters to estimate, the range and the variance (sill). The results are shown in Table 1 for a sample size of 30 and in Table 2 for a sample size of 90.

The conclusions for the range are:

- The best estimator (in the sense of minimum RMSE) is ML for a sample size of 30 and MLCV2 for a sample size of 90 (although ML is second and close to MLCV2).
- The ML methods have a negative bias.
- Among the LS estimators GLS is the best followed by WLS2.
- There is no information lost in binning as GLS and WLS2 perform better than methods that use the semivariogram cloud (OLS and WLS1).
- WLS1 performs better than OLS, thus it is better not to use the semivariogram values for long distances (high uncertainty).

Table 1

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget variance. Sample size is 30.

<i>n</i> =30	Mean of sill (target=10)	Variance of sill	Mean of range (target=6)	Variance of range
OLS	11.04	27.49	8.12	64.08
WLS1	10.75	27.78	6.92	42.97
WLS2	11.08	31.25	3.84	23.29
GLS	10.71	32.76	6.58	41.02
ML	9.78	26.25	4.43	9.74
CV1	13.17	15.24	8.70	41.99
MLCV1	11.04	30.79	6.48	35.39
CV2	13.10	22.29	9.25	40.72
MLCV2	10.73	28.39	5.69	21.81

<i>n</i> =30	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	1.04	5.34	2.12	8.28
WLS1	0.75	5.32	0.92	6.62
WLS2	1.08	5.69	-2.16	5.28
GLS	0.71	5.76	0.58	6.43
ML	-0.22	5.12	-1.57	3.49
CV1	3.17	5.02	2.70	7.02
MLCV1	1.04	5.64	0.48	5.96
CV2	3.10	5.64	3.25	7.16
MLCV2	0.73	5.37	-0.31	4.68

Table 3

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with nugget variance equal to 50% of the total variance. Sample size is 30. This table shows the sill and range parameters.

<i>n</i> =30	Mean of sill (target=10)	Variance of sill	Mean of range (target=6)	Variance of range
OLS	11.17	25.36	13.17	96.64
WLS1	11.34	30.38	13.13	93.33
WLS2	11.11	28.77	10.78	91.21
GLS	10.79	26.76	10.46	88.84
ML	9.45	12.36	4.06	19.63
CV1	14.52	25.15	13.43	47.70
MLCV1	10.89	23.23	10.51	98.05
CV2	12.40	22.30	13.94	41.31
MLCV2	10.44	17.84	9.28	87.56

<i>n</i> =30	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	1.17	5.17	7.17	12.16
WLS1	1.34	5.67	7.13	12.00
WLS2	1.11	5.47	4.78	10.68
GLS	0.79	5.23	4.46	10.42
ML	-0.55	3.55	-1.94	4.83
CV1	4.52	6.75	7.43	10.14
MLCV1	0.89	4.90	4.51	10.88
CV2	2.40	5.29	7.94	10.21
MLCV2	0.44	4.24	3.28	9.91

Table 2

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget variance. Sample size is 90.

<i>n</i> =90	Mean of sill	Variance of sill	Mean of range	Variance of range
OLS	10.10	28.19	8.55	63.74
WLS1	9.26	28.23	5.66	27.84
WLS2	9.28	29.54	2.65	8.54
GLS	8.61	27.97	5.86	41.10
ML	9.21	27.38	3.91	6.64
CV1	12.38	17.45	6.78	24.59
MLCV1	9.46	30.52	4.44	17.76
CV2	12.77	17.35	7.83	33.74
MLCV2	9.68	28.63	4.29	7.74

<i>n</i> =90	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	0.10	5.31	2.55	8.38
WLS1	-0.74	5.36	-0.34	5.28
WLS2	-0.72	5.48	-3.35	4.44
GLS	-1.39	5.46	-0.14	6.41
ML	-0.79	5.29	-2.09	3.31
CV1	2.38	4.80	0.78	5.02
MLCV1	-0.54	5.55	-1.56	4.49
CV2	2.77	5.00	1.83	6.09
MLCV2	-0.32	5.36	-1.71	3.26

Table 4

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with nugget variance equal to 50% of the total variance. Sample size is 30. This Table shows the nugget variance and partial sill parameters.

<i>n</i> =30	Mean of nugget (target=5)	Variance of nugget	Mean of partial sill (target=5)	Variance of partial sill
OLS	4.32	10.93	6.85	27.72
WLS1	4.51	12.37	6.83	28.52
WLS2	4.38	15.67	6.73	27.37
GLS	3.91	13.08	6.88	25.86
ML	3.87	10.71	5.58	20.88
CV1	4.93	7.70	9.59	25.64
MLCV1	4.23	12.15	6.66	30.08
CV2	4.51	11.51	7.89	24.15
MLCV2	4.40	11.08	6.04	23.13

<i>n</i> =30	Bias of nugget	RMSE of nugget	Bias of partial sill	RMSE of partial sill
OLS	-0.68	3.37	1.85	5.58
WLS1	-0.49	3.55	1.83	5.64
WLS2	-0.62	4.00	1.73	5.51
GLS	-1.09	3.77	1.88	5.42
ML	-1.13	3.46	0.58	4.60
CV1	-0.07	2.77	4.59	6.83
MLCV1	-0.77	3.57	1.66	5.73
CV2	-0.49	3.42	2.89	5.70
MLCV2	-0.60	3.38	1.04	4.92

- CV methods are improved by including the global measure of overall accuracy, i.e. converting them to MLCV methods.
- Orthonormal residuals provide better results than ordinary residuals.

Conclusions for the sill:

- All methods have very similar RMSE.

4.1.2. Non-Gaussian RF with nugget variance

In this case there are three parameters to estimate, the nugget variance, the partial sill and the range. The variance (total sill) is estimated as the sum of nugget variance and partial sill. The results are shown in Tables 3 and 4 for a sample size of 30. Table 3

shows the results for the sill and range parameters and Table 4 shows the results for nugget variance and partial sill parameters. Tables 5 and 6 show the results for a sample size of 90. Table 5 shows the results for the sill and range parameters and Table 6 shows the results for nugget variance and partial sill parameters.

Conclusions for the range:

- The best estimator (in the sense of minimum RMSE) is ML for a sample size of 30 followed closely by GLS. The remaining methods have significantly higher RMSE.
- ML has a negative bias.
- There is no information lost in binning as GLS and WLS2 perform better than methods that use the semivariogram cloud (OLS and WLS1).

Table 5

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with nugget variance equal to 50% of the total variance. Sample size is 90. This Table shows the sill and range parameters.

<i>n</i> =90	Mean of sill (target=10)	Variance of sill	Mean of range (target=6)	Variance of range
OLS	12.07	15.26	15.82	86.02
WLS1	11.38	13.81	12.92	72.87
WLS2	11.19	13.91	11.34	81.90
GLS	11.04	14.79	9.68	81.23
ML	10.22	8.61	5.56	28.62
CV1	14.82	18.64	13.3	35.73
MLCV1	12.02	28.94	9.38	90.49
CV2	13.43	21.12	14.04	47.23
MLCV2	10.44	17.84	9.28	87.56
<i>n</i> =90	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	2.07	4.42	9.82	13.50
WLS1	1.38	3.96	6.92	10.98
WLS2	1.19	3.91	5.34	10.50
GLS	1.04	3.98	3.68	9.73
ML	0.22	2.94	-0.44	5.36
CV1	4.82	6.47	7.30	9.43
MLCV1	2.02	5.74	3.38	10.09
CV2	3.43	5.73	8.04	10.57
MLCV2	1.12	3.69	3.30	9.72

Table 6

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with nugget variance equal to 50% of the total variance. Sample size is 90. This Table shows the nugget variance and partial sill parameters.

<i>n</i> =90	Mean of nugget (target=5)	Variance of nugget	Mean of partial sill (target=5)	Variance of partial sill
OLS	5.03	7.24	7.04	19.67
WLS1	5.31	5.95	6.07	18.72
WLS2	4.92	7.43	6.27	16.53
GLS	4.6	7.24	6.44	17.12
ML	4.13	7.09	6.09	13.58
CV1	5.55	5.90	9.27	26.09
MLCV1	3.99	7.27	8.03	31.26
CV2	4.65	7.70	8.78	28.49
MLCV2	3.77	11.49	7.35	20.08
<i>n</i> =90	Bias of nugget	RMSE of nugget	Bias of partial sill	RMSE of partial sill
OLS	0.03	2.69	2.04	4.88
WLS1	0.31	2.46	1.07	4.45
WLS2	-0.08	2.72	1.27	4.26
GLS	-0.40	2.72	1.44	4.38
ML	-0.87	2.80	1.09	3.84
CV1	0.55	2.49	4.27	6.65
MLCV1	-1.01	2.87	3.03	6.36
CV2	-0.35	2.79	3.78	6.54
MLCV2	-1.23	3.60	2.35	5.06

- Orthonormal residuals provide better results than ordinary residuals but the difference is less than when there is no nugget.
- The RMSE is higher than for the case with no nugget (Table 1) and for all the estimators. That is, the presence of noise in the model increases the noise in the estimates and increases the uncertainty of the estimators.

Conclusions for the variance (total sill):

- The RMSE is similar for all estimators and similar to the no nugget case except for the GLS, which performs poorly.

Conclusions for the nugget variance:

- CV1 gives the smallest RMSE. The other methods have similar performances except GLS, which is the worst performer.

Conclusions for the partial sill:

- ML gives the smallest RMSE. The other methods have similar performances.

4.1.3. Non-Gaussian RF without nugget-variance but where the nugget variance parameter is estimated

This example shows the consequence of estimating the nugget variance to be zero when in fact the RF has no nugget.

The results are shown in Tables 7 and 8 for a sample size of 30. Table 7 shows the results for the sill and range parameters and Table 8 shows the results for nugget variance and partial sill parameters. Tables 9 and 10 show the results for a sample of size 90. Table 9 shows the results for the sill and range parameters and Table 10 shows the results for nugget variance and partial sill parameters.

Conclusions for the range:

- The RMSE increases for all the methods. Thus if, from physical principles or expert knowledge, it is known that there is no nugget variance then it is better not to include the parameter in the estimation process.
- ML is the best performer in the sense that its RMSE is a little higher than when the nugget parameter is not included in the inference while the RMSE for other estimators, including MLCV2, is double the value achieved when the nugget parameter is not included in the inference.

Conclusions for the variance:

- The RMSE for all methods except MLCV2 is increased but the increase in RMSE for ML is the smallest.

Conclusions for the nugget variance:

- All methods estimate a nugget greater than zero. The smallest bias (i.e. value closest to zero) is for ML for *n*=30 and MLCV1 for *n*=90.

Table 7

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget variance. Sample size is 30. This Table shows the sill and range parameters.

<i>n</i> =30	Mean of sill (target=10)	Variance of sill	Mean of range (target=6)	Variance of range
OLS	12.52	52.03	11.74	87.15
WLS1	12.45	60.24	10.62	69.27
WLS2	12.17	56.28	9.56	58.94
GLS	11.83	49.02	8.66	57.18
ML	10.04	31.55	5.93	20.74
CV1	12.98	43.58	14.93	42.48
MLCV1	12.64	36.41	11.08	70.91
CV2	12.30	49.13	13.9	50.39
MLCV2	10.61	19.25	10.03	89.98
<i>n</i> =30	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	2.52	7.64	5.74	10.95
WLS1	2.45	8.13	4.62	9.51
WLS2	2.17	7.81	3.56	8.46
GLS	1.83	7.23	2.66	8.01
ML	0.04	5.61	-0.07	4.55
CV1	2.98	7.24	8.93	11.05
MLCV1	2.64	6.58	5.08	9.83
CV2	2.3	7.37	7.9	10.62
MLCV2	0.61	4.43	4.03	10.30

Table 8

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget. Sample size is 30. This Table shows the nugget variance and partial sill parameters.

<i>n</i> =30	Mean of nugget (target=0)	Variance of nugget	Mean of partial sill (target=10)	Variance of partial sill
OLS	2.36	13.01	10.16	30.39
WLS1	2.38	14.75	10.07	33.74
WLS2	2.07	12.70	10.10	33.01
GLS	1.69	10.01	10.14	32.70
ML	1.31	5.61	8.73	28.25
CV1	4.87	19.73	8.11	23.67
MLCV1	1.22	4.37	11.42	35.22
CV2	4.92	21.81	7.38	25.85
MLCV2	3.91	13.06	6.70	24.39

<i>n</i> =30	Bias of nugget	RMSE of nugget	Bias of partial sill	RMSE of partial sill
OLS	2.36	4.31	0.16	5.51
WLS1	2.38	4.51	0.07	5.80
WLS2	2.07	4.12	0.10	5.74
GLS	1.69	3.58	0.14	5.72
ML	1.31	2.70	-1.27	5.46
CV1	4.87	6.59	-1.89	5.22
MLCV1	1.22	2.42	1.42	6.10
CV2	4.92	6.78	-2.62	5.72
MLCV2	3.91	5.32	-3.30	5.94

Table 9

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget variance. Sample size is 90. This Table shows the sill and range parameters.

<i>n</i> =90	Mean of sill (target=10)	Variance of sill	Mean of range (target=6)	Variance of range
OLS	10.89	40.35	13.61	95.63
WLS1	10.19	40.35	10.96	72.97
WLS2	9.94	37.91	9.29	55.28
GLS	9.78	34.01	8.47	57.98
ML	9.29	27.46	5.82	17.30
CV1	11.52	29.09	14.63	34.73
MLCV1	11.39	38.25	11.33	70.12
CV2	11.14	34.12	13.02	40.58
MLCV2	11.12	12.42	9.30	83.59

<i>n</i> =90	Bias of sill	RMSE of sill	Bias of range	RMSE of range
OLS	0.89	6.41	7.61	12.39
WLS1	0.19	6.35	4.96	9.87
WLS2	-0.06	6.15	3.29	8.13
GLS	-0.22	5.83	2.47	8.00
ML	-0.71	5.28	-0.18	4.16
CV1	1.52	5.60	8.63	10.45
MLCV1	1.39	6.34	5.33	9.92
CV2	1.14	5.95	7.02	9.47
MLCV2	1.12	3.69	3.30	9.72

Conclusions for the partial sill:

- The partial sill, which is equal to the sill or total variance because the nugget variance of the RF is zero, is estimated with smaller RMSE than the variance because the variance is estimated as the sum of partial sill and nugget variance.

4.2. Joint estimation of several parameters

In order to evaluate the performance of the methods, the RMSE is accumulated and the methods are ranked according to their scores. The results are shown in Table 11. In increasing order of total RMSE, the ranked methods are: ML (259), MLCV2 (340),

Table 10

Statistics of the sampling distribution of the estimates for a non-Gaussian RF and a practical range/side length ratio of 0.4. RF with no nugget. Sample size is 90.

<i>n</i> =90	Mean of nugget (target=0)	Variance of nugget	Mean of partial sill (target=10)	Variance of partial sill
OLS	1.77	4.41	9.12	32.12
WLS1	1.90	3.79	8.29	33.40
WLS2	1.86	3.36	8.08	29.91
GLS	1.55	3.66	8.23	29.01
ML	1.14	2.06	8.15	26.40
CV1	3.09	7.90	8.43	28.92
MLCV1	1.33	2.94	10.06	35.73
CV2	4.18	18.42	6.96	21.11
MLCV2	3.77	11.49	7.35	20.08

<i>n</i> =90	Bias of nugget	RMSE of nugget	Bias of partial sill	RMSE of partial sill
OLS	-3.23	3.85	4.12	7.00
WLS1	-3.10	3.66	3.29	6.65
WLS2	-3.14	3.63	3.08	6.27
GLS	-3.45	3.94	3.23	6.28
ML	-3.86	4.11	3.15	6.02
CV1	-1.91	3.39	3.43	6.37
MLCV1	-3.67	4.05	5.06	7.83
CV2	-0.82	4.37	1.96	4.99
MLCV2	-1.23	3.60	2.35	5.06

Table 11

Scoring of accumulated RMSE (in Tables 1–10) of the estimators for each parameter and for the total model.

Accumulated RMSE	RMSE of nugget	RMSE of partial sill	RMSE of variance	RMSE of range	Total RMSE
OLS	13	22	34	66	416
WLS1	13	22	35	54	379
WLS2	13	21	35	48	357
GLS	13	21	34	49	366
ML	11	19	28	26	259
CV1	16	24	36	53	389
MLCV1	11	24	35	51	370
CV2	19	23	35	54	396
MLCV2	18	21	27	48	340

WLS2 (357), GLS (366), MLCV1 (370), WLS1 (379), CV1 (389), CV2 (396), and OLS (416). For the range the best is ML (26) with almost half the RMSE of any other method. MLCV2 is second (48). For the nugget variance, ML (11) and MLCV1 (11) are the best. For the partial sill ML is the best (19) closely followed by the three methods (21). Finally, for the variance (sill) MLCV2 is the best (27) closely followed by ML (28).

4.3. Assessment of the uncertainty of the estimates

We can compare the uncertainty evaluations by comparing the coverage and the width of the intervals. However, it should be borne in mind that we have chosen to use the worst-case scenario in which the RF is highly non-Gaussian, there is no transformation of the data and no transformation of the estimates. On the other hand all the uncertainty evaluations are parametric with Gaussian or chi-square distribution assumptions for the given statistics, i.e., Eqs. (14), (15), (18) and (19).

By way of example, the results of Eqs. (14) and (15) for GLS and ML are given in Table 12. Because the assumptions are not correct there is no correspondence between the actual coverage and the nominal coverage under the Gaussian assumption. This requires further work, for example, implementing non-parametric approaches like a bootstrap procedure.

Table 12

Coverage of random intervals for GLS and ML. For a Gaussian distribution the coverage of estimate $\pm 1SE$ is of 68% and estimate $\pm 2SE$ has a coverage of 95%.

Coverages	Variance estimate $\pm 1SE$	Variance estimate $\pm 2SE$	Range estimate $\pm 1SE$	Range estimate $\pm 2SE$
GLS 30	55	72	65	78
GLS 90	36	63	41	57
ML 30	43	68	95	98
ML 90	21	48	75	88

5. Conclusions

The primary purpose of this paper is to compare methods of inferring semivariogram model parameters. The study is limited to small samples from a two-dimensional, second-order stationary non-Gaussian random function with an isotropic semivariogram model with and without a nugget effect.

The OLS estimator using the full semivariogram cloud is the worst performer among the set of estimators that have been compared. The estimator may be improved in two ways: (1) by using WLS1 with zero weights applied to those data pairs in the semivariogram cloud that are separated by significant distances. Uncertainty of the semivariogram values increases with distance and beyond a certain limit there is no value in including them in the estimator; (2) by using WLS2 and using only semivariogram lags up to a given distance (for example one third of the study area or three times the expected range). GLS gives similar results to WLS2. This may be because we have used the worst-case scenario with a highly skewed non-Gaussian RF from which the original distribution cannot be recovered.

Among the cross-validation estimators those that use the global measure of accuracy of Kitanidis (1991) have proved superior to the others. Estimators that use the global measure are equivalent to the maximum likelihood cross-validation method. In addition, the orthonormal residuals give better results than classical residuals. Thus MLCV2 is the best among the cross-validation estimators and the second best among all the estimators tested.

ML is the best estimator even when the data are non-Gaussian. ML is a good measure of the goodness of fit of a semivariogram model to the experimental data.

If there is no reason to believe that a nugget variance is present in the data, the results can be improved by not including this parameter in the inference, i.e., implicitly taking it to be zero.

Finally, further research is needed on the estimation of uncertainty measures with a given probabilistic coverage in the worst-case scenario.

Acknowledgments

This work was supported by research projects CGL2010-15498 and CGL2010-17629 from the Ministerio de Economía y Competitividad of Spain and by the Australian Research Council through Discovery Project Grant no. DP110104766.

References

Bard, Y., 1974. *Nonlinear Parameter Estimation*. John Wiley and Sons, New York 341 p.

Bogaert, P., Russo, D., 1999. Optimal spatial sampling design for the estimation of the variogram based on a least squares approach. *Water Resources Research* 35 (4), 1275–1289.

Burrough, P.A., 2001. GIS and geostatistics: essential partners for spatial analysis. *Environmental and Ecological Statistics* 8, 361–377.

Caers, J., 2005. *Petroleum Geostatistics*. Society of Petroleum Engineers, Richardson, TX 88 p.

Chica-Olmo, M., Abarca-Hernández, F., 2000. Computing geostatistical image texture for remotely sensed data classification. *Computers & Geosciences* 26, 373–383.

Chilès, J.P., Delfiner, P., 2012. *Geostatistics. Modeling Spatial Uncertainty*, Second edition John Wiley & Sons, Hoboken, New Jersey 702 p.

Cressie, N., 1985. Fitting variogram models by weighted least squares. *Mathematical Geology* 17 (5), 563–586.

Curran, P.J., Atkinson, P.M., 1998. Geostatistics and remote sensing. *Progress in Physical Geography* 22 (1), 61–78.

Deutsch, C.V., 2002. *Geostatistical Reservoir Modeling*. Oxford University Press, New York 384 p.

Deutsch, C.V., Journel, A.G., 1998. *GSLIB. Geostatistical Software Library and Users Guide*, Second edition Oxford University Press 384 p.

Diggle, P.J., Ribeiro Jr, P.J., 2007. *Model Based Geostatistics*. Springer-Verlag, New York 228 p.

Diodato, N., Ceccarelli, M., 2005. Interpolation processes using multivariate geostatistics for mapping of climatological precipitation mean in the Sannio Mountains (southern Italy). *Earth Surface Processes and Landforms* 30 (3), 259–268.

Dowd, P.A., 1983. The variogram and kriging: robust and resistant estimators. In: Verly, G., et al. (Eds.), *Geostatistics for Natural Resource Characterization Part 1: NATO A.S.I. Series C: Mathematical and Physical Sciences*, vol. 122. D. Reidel Publishing Co., Dordrecht, Netherlands, pp. 91–106.

Genton, M.G., 1998. Highly robust variogram estimation. *Mathematical Geology* 30 (2), 213–221.

Griffith, D.A., Paelinck, J.H.P., 2011. *Non-Standard Spatial Statistics and Spatial Econometrics*. Springer, Berlin 262 p.

Goovaerts, P., 1997. *Geostatistics for Natural Resources Evaluation*. Oxford University Press, New York 463 p.

Gringarten, E., Deutsch, C.V., 2001. Teacher's aide variogram interpretation and modeling. *Mathematical Geology* 33 (4), 507–534.

Journel, A.G., Rossi, M.E., 1989. When do we need a trend model in kriging. *Mathematical Geology* 21 (7), 715–739.

Kalbfleisch, J.G., 1979. *Probability and Statistical Inference II*. Springer-Verlag, New York.

Kelsall, J., Wakefield, J., 2002. Modeling spatial variation in disease risk: a geostatistical approach. *Journal of the American Statistical Association* 97 (459), 692–701.

Kitanidis, P.K., 1983. Statistical estimation of polynomial generalized covariance functions and hydrologic applications. *Water Resources Research* 19 (4), 909–921.

Kitanidis, P.K., 1991. Orthonormal residuals in geostatistics: model criticism and parameter estimation. *Mathematical Geology* 23 (5), 741–757.

Kitanidis, P.K., 1997. *Introduction to Geostatistics. Applications to Hydrogeology*. Cambridge University Press, Cambridge 249 p.

Kitanidis, P.K., Lane, R.W., 1985. Maximum likelihood parameter estimation of hydrologic spatial processes by the Gaussian-Newton method. *Journal of Hydrology* 79 (1–2), 53–71.

Krajewski, W.F., Duffy, C.J., 1988. Estimation of correlation structure for a homogeneous isotropic random field: a simulation study. *Computers & Geosciences* 14 (1), 113–122.

Lebel, T., Bastin, G., 1985. Variogram identification by mean squared interpolation error method with application to hydrologic fields. *Journal of Hydrology* 77 (1–4), 31–56.

Liebhold, A.M., Rossi, R.E., Kemp, W.P., 1993. Geostatistics and geographic information systems in applied insect ecology. *Annual Review of Entomology* 38, 303–327.

Mardia, K.V., Marshall, R.J., 1984. Maximum likelihood estimation of models for residual covariance in spatial regression. *Biometrika* 71 (1), 135–146.

Matheron, G., 1963. Principles of geostatistics. *Economic Geology* 58, 1246–1266.

Menke, W., 1984. *Geophysical Data Analysis: Discrete Inverse Theory*. Academic Press, San Diego, CA 285 p.

McCullagh, P., Nelder, J.A., 1989. *Generalized Linear Models*, 2nd ed. Chapman and Hall, Cambridge, Massachusetts.

O'Brien, J.J., Griffiths, J.F., 1965. The rank correlation coefficient as an indicator of the product-moment correlation coefficient for small samples (10–100). *Journal of Geophysical Research* 70 (8), 1995–1998.

Olea, R., Pardo-Igúzquiza, E., 2011. Generalized bootstrap method for assessing of uncertainty in semivariogram inference. *Mathematical Geosciences* 43 (2), 203–228.

Pardo-Igúzquiza, E., 1998. Maximum likelihood estimation of spatial covariance parameters. *Mathematical Geology* 30 (1), 95–108.

Pardo-Igúzquiza, E., 1999a. Bayesian inference of spatial covariance parameters. *Mathematical Geology* 31 (1), 47–65.

Pardo-Igúzquiza, E., 1999b. VAREFIT: a fortran-77 program for fitting variogram models by weighted least squares. *Computer & Geosciences* 25 (3), 251–261.

Pardo-Igúzquiza, E., Chica-Olmo, M., Luque-Espinar, J.A., García-Soldado, M.J., 2009. Using semivariogram parameter uncertainty in hydrogeological applications. *Groundwater* 47 (1), 25–34.

Pardo-Igúzquiza, E., Dowd, P.A., 2001. Variance-covariance matrix of the experimental variogram: assessing variogram uncertainty. *Mathematical Geology* 33 (4), 397–419.

Pardo-Igúzquiza, E., Dowd, P.A., 2003. Assessment of the uncertainty of spatial covariance parameters of soil properties and its use in applications. *Soil Science* 168 (11), 769–782.

- Pardo-Igúzquiza, E., Dowd, P.A., 2005. Empirical maximum likelihood kriging: the general case. *Mathematical Geology* 37 (5), 477–492.
- Rice, J.A., 1995. *Mathematical Statistics and Data Analysis*, Second edition Duxbury Press, Belmont, California 602 p.
- Russo, D., Jury, W.A., 1987. A theoretical study of the estimation of the correlation scale in spatially variable fields: 1. Stationary fields. *Water Resources Research* 23 (7), 1257–1268.
- Samper, F.J., Neuman, S.P., 1989. Estimation of spatial covariance structures by adjoint state maximum likelihood cross-validation: 1. Theory. *Water Resources Research* 25 (3), 351–362.
- Sampson, P.D., Guttorp, P., 1992. Nonparametric estimation of nonstationary spatial covariance structure. *Journal of the American Statistical Association* 87 (417), 108–119.
- Solow, A.R., 1985. Bootstrapping correlated data. *Mathematical Geology* 17 (7), 769–775.
- Spiegel, M.R., Stephens, L.J., 2008. *Statistics*, Fourth edition McGraw-Hill Professional, New York 577 p.
- Strebelle, S., 2002. Conditional simulation of complex geological structures using multiple-point statistics. *Mathematical Geology* 34, 1–22.
- Visser, A., Dubus, I., Broers, H.P., Brouyère, S., Korcz, M., Orban, P., Goderniaux, P., Batlle-Aguilar, J., Surdyk, N., Amraoui, N., Job, H., Pinault, J.L., Bierkens, M., 2009. Comparison of methods for the detection and extrapolation of trends in groundwater quality. *Journal of Environmental Monitoring* 11, 2030–2043.
- Ye, M., Meyer, P.D., Neuman, S.P., 2008. On model selection criteria in multimodel analysis. *Water Resources Research* 44 (W03428), 1–12.
- Zimmerman, D.L., Zimmerman, M.B., 1991. A comparison of spatial semivariogram estimators and corresponding ordinary kriging predictors. *Technometrics* 33 (1), 77–91.